## Approximate solution of bidimensional momentum equations for inviscid atmospheric flows

Marco A. Núñez, Gerardo Ramírez R.

Depto. Física, Universidad Autónoma Metropolitana Iztapalapa, AP 55-534, CP 09340, D.F., México. e-mail: manp@xanum.uam.mx

It is generally accepted [1] that the movement of large scale atmospheric flows is governed by the bidimensional Euler equations

$$\frac{du}{dt} - \frac{uv\tan\phi}{r_e} - 2\Omega v\sin\phi = -\frac{1}{\rho}\frac{\partial p}{\partial x}$$

$$\frac{dv}{dt} + \frac{u^2\tan\phi}{r_e} + 2\Omega u\sin\phi = -\frac{1}{\rho}\frac{\partial p}{\partial y}$$
(1)

where  $\frac{d}{dt} = \frac{\partial}{\partial t} + u \frac{\partial}{\partial x} + v \frac{\partial}{\partial y}$ . In this work we give an approximate analytic solution  $u_a$ ,  $v_a$  of Eqs. (1), which was obtained by a perturbation method [2].

The accuracy of  $u_a$ ,  $v_a$  is estimated by computing the solutions u, v corresponding to the inertial trajectories of an atmospheric particle, since in this case u, v can be computed analytically [3]. There are several regimens of qualitatively different inertial trajectories, regimens that can be classified with the aid of the Hamiltonian formulation of (1) [3]. The results reported in this work show that  $u_a$ ,  $v_a$  are close to u, v as t goes from 0 to  $t_{\max} \sim 18$  hrs independently of the regimen of movement. The criterion to estimate  $t_{\max}$  is the following: if  $\mathbf{r}_a(t)$  and  $\mathbf{r}(t)$  are the inertial trajectories obtained from  $\{u_a, v_a\}$  and  $\{u, v\}$ , respectively, we set  $d(\mathbf{r}_a, \mathbf{r}) = \|\mathbf{r}_a(t) - \mathbf{r}(t)\|^{1/2}$  and define  $t_{\max} = \inf\{t : d(\mathbf{r}_a, \mathbf{r}) = 10 \text{ km}\}$ . The reliability time interval  $[0, t_{\max}]$  increases as the latitude of the movement also does, in such a way that for latitudes from  $55^{\circ}$  to  $85^{\circ}$  the approximations  $u_a$ ,  $v_a$  are close to u, v for  $t \in [0, t_{\max} \sim 35 \text{ hrs}]$ . Additional graphs of t v.s.  $\{u_a, v_a\}$ ,  $\{u, v\}$  confirm the closness between  $u_a, v_a$  and u, v. These results suggest that the approximations  $u_a$ ,  $v_a$  can be used for operational and theoretical purposes.

References.

J. R. Holton, An introduction to dynamic meteorology, 3rd. ed. (Academic Press, San Diego, 1992).
 J. Pedloski, Geophysical Fluid Dynamics (Springer, New York, 1986).

[2] K. Kevorkian and J. D. Cole, Multiple Scale and Singular Perturbation Methods (Springer, New York, 1996). M. Van Dike, Perturbation methods in fluid mechanics (Academic Press, London, 1964).
[3] N. Paldor and P. D. Killworth, Inertial trajectories on a rotating Earth, J. Atmos. Sci. 45, 4013-4019 (1988). N. Paldor and E. Boss Chaotic Trajectories of Tidally Perturbed Inertial Oscillations, J. Atmos. Sci. 49, 2306-2318 (1992).